

Five years of gravity measurement at Onsala Space Observatory: The absolute scale

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The gravimetry lab

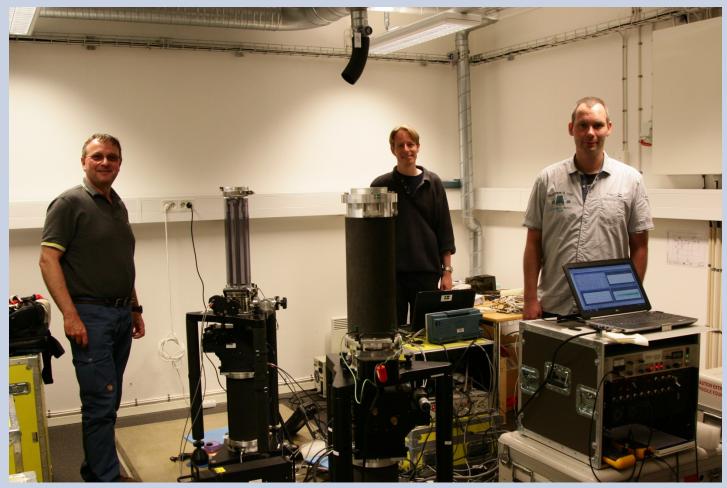
at Onsala Space Observatory



Tide gauges: Bubble Radar GNSS



The parallel campaign May 2014



Ludger Timmen

FG5X 220

FG5 233 Andreas Engfeldt

Manuel Schilling

4 Scherneck Engfeldt Olsson Timmen

Content

• The limitations of the geophysical models in g-software

- Atmosphere and seasonal perturbations
- Ocean loading: nontidal, non-static, non-stationary

• FG5 absolute gravimeters

Drift on the Project time-scale

• The limitations of SCG-measurements

– Drift

• The Atmacs atmospheric attraction and loading model

Non-static effects from short to long time-scale

• Reduction of AG-measurements by SCG-data

- At the Drop-level

AG reduction of observations by models

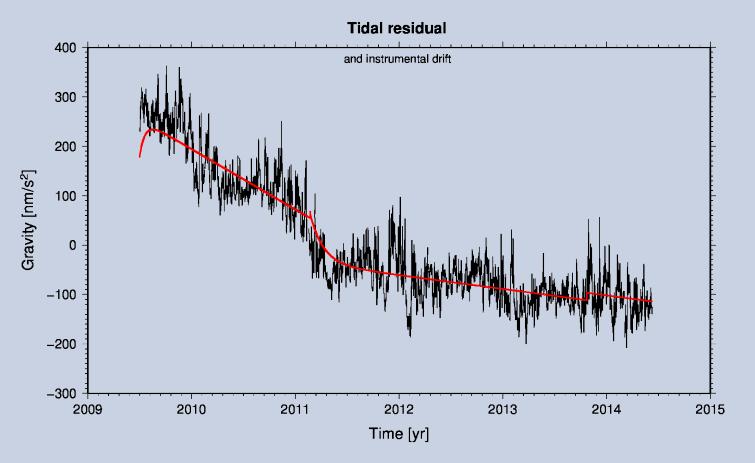
- The standard method, geophyscial models:
 - Polar motion, tides, ocean loading, atmosphere, ...
- AG ... by direct application of SCG-observations
 - Sensitivity that AG and SCG have in common:
 - Polar motion, tides, ocean loading, atmosphere, hydrosphere
- Data, no model:
 - No need to assume stationarity,
 - SCG drift needs a model though

First I will show that we "understand" the SCG record down to the 2 nm/s² level

- To first order: a complex, stationary model
 - Tides, polar motion, sea level, atmosphere (two shades of Atmacs)
- Additional dynamics "explained" by stochastic models (Wiener filtering) = "non-static"
- Tides: Empirical tide coefficients in which we can identify known effects (tides, ocean tide loading) but also find effects yet to be identified

- Cannot address AG instrumental limitations
- Devote efforts to SCG instrumental limitations, primarily DRIFT
- The challenges of a wide-band signal
 - 1-second sampling at the front-end of processing
 - 1-hour averages, a compromise
 - atmosphere model Atmacs (3 h) => spectral fill-in local barometer (not so simple as it may sound)
 - sea-level data Ringhals (1 h) => decimate OSO tide gauge data
 - 1-day averages (nice data, not too useful) => testing Atmacs at long periods
- By the way:

Atmacs: <u>Atmospheric Attraction Computation Service</u> http://atmacs.bkg.bund.de



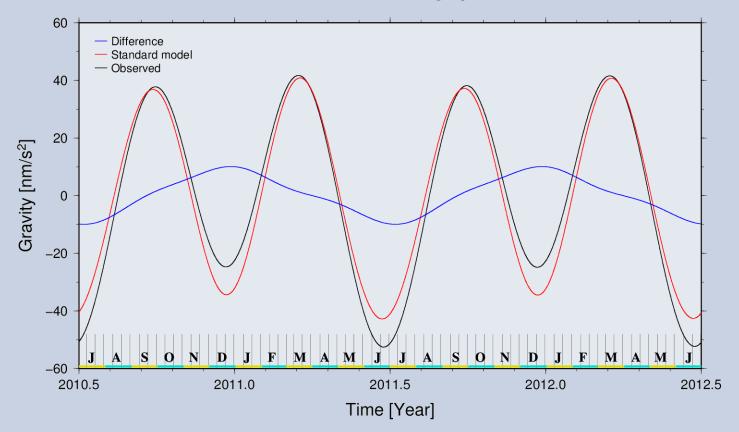
Is this difference, Tide-residual minus drift, a reliable prediction for AG?

Notice:

long-term excursions are neither exactly annual (S_a tide coeff. took this effect away) nor related to Polar Motion (that's in the model too).

Long-period, annual, seasonal

Excess solar long-period



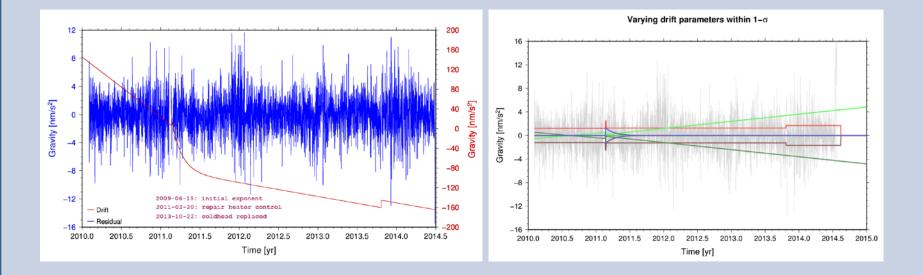
See the poster on this. A short summary ...

Some findings

- The annual tide at OSO is 2 × the solar gravity
- This observation is ≈ independent of using either 1-hour data with spectral whitening or 1-day data without
- ... while the admittance coefficients for Atmacs regional+local and global change substantially (a matter of best fit at daily periods vs at very-long periods.
- The perturbation peaks low-to-high: early January to mid June. See the poster
- Presently we don't know the origin (while we are ignorant of ground water level, hydrology, biomass...

(a ground water monitor below the lab is on the way))

• The SCG will "know" (beware an unknown instrumental effect – hmm)



Assuming that we can determine drift components (slopes, biases, exponentials) at the 5 nm/s² level, we could try to take the rest of the SCG observations at face value.

The residual RMS in the extended solution:2.27 nm/s²In the standard solution (no Wiener filters):7.05 nm/s²

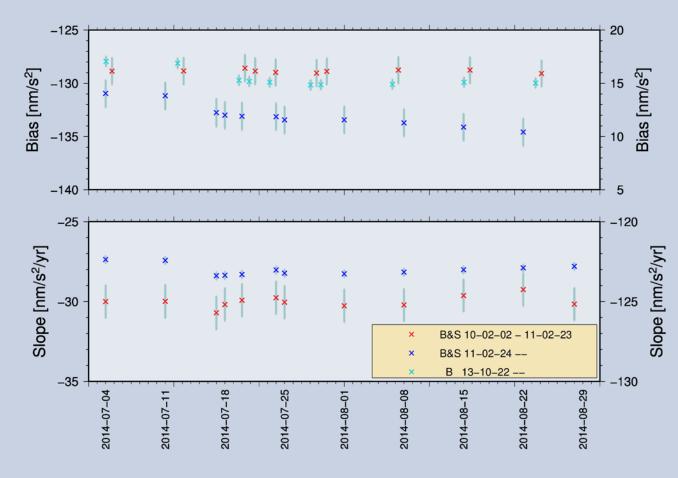
The residual RMS in the extended solution is

2.27 nm/s² !!!

from 40,000 hours of data

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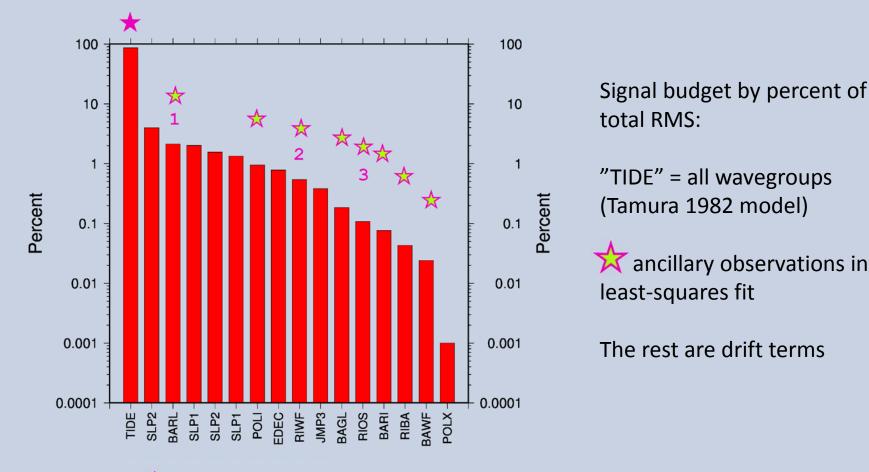
Drift parameters - robustness



We must keep an eye on the drift parameters. So far they spread wider than the standard deviation. Since May 2014 we produce a standard solution every week. (In July the setup was changed):

Atmacs R+L Atmacs Glob Ringhals+OSO Sea level Polar motion 63 tide wave groups 6 drift parameters (5 shown here)

Recent change B&S: 2014-08-28 affects biases



Comments: 4 - Regional+Local atmosphere model from Atmacs,

- not surprisingly most important after tides
- 2 Tide gauge (mostly from Ringhals, lately from OSO) Wiener-filtered "dynamic admittance" (although a stationary concept)
 - more important than:
- ☆ 3 Tide gauge: the static admittance

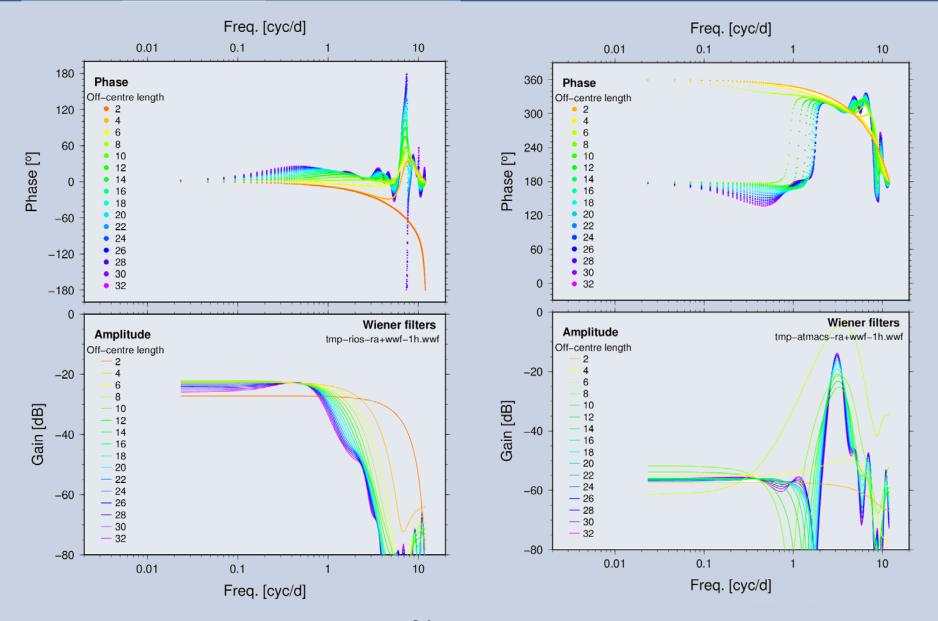
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Now I have to explain...

Cross-correlations with g-residual After a least-squares fit with single admittance coefficients 2.0 for Atmacs-R+L and tide 1.5 gauge, RIOS-BARL GR := q - tides - a AT - t TG1.0 off-centre cross correlation is 0.5 Correlation *3 found between these time BARL 0.0 series: $\{GR \mid AT\} \{GR \mid TG\}$ -0.5RIOS and between the pair 80.50.683.55.083.55.053.683.693.673.676.826.000.0765.570.0766.6520.000.6765.65 -1.0{ $TG - \alpha AT \mid AT - \tau TG$ } -1.5 (over-simplified, sorry) -2.0-100 100 -200 0 200 <= this information is the Lag [h] basis for the Wiener filters:

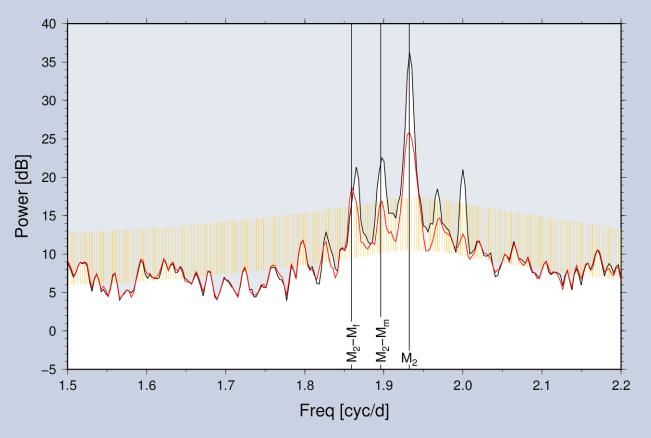
 $Wf = \mathcal{F}^{-1}\{Cohc \cdot XY^*/|Y|^2\} \cdot w$

with a window for tapering off at large lag (noise suppression) and the coherence spectrum as a multiplier (selecting common features)



{ GR | TG } tide gauge <- Wiener filter gain spectra -> { GR | AT } Atmacs R&L

Nonstationary M₂ found in the power spectrum



Ringhals tide gauge:

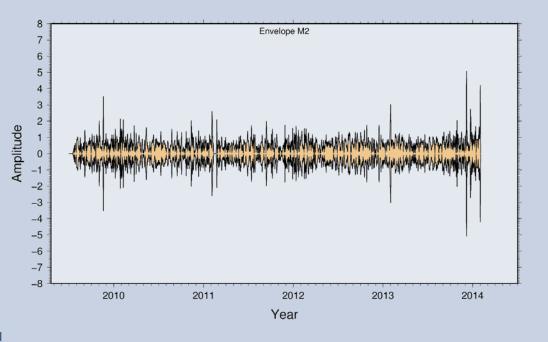
Black: Unreduced for tides.

Red:

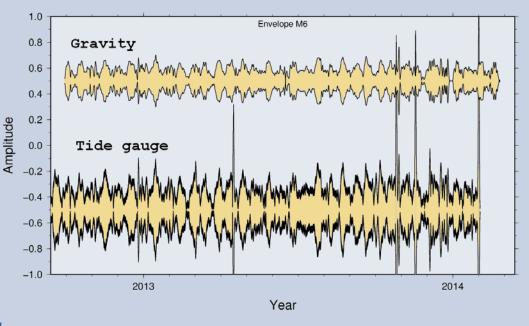
After reduction of the tide wave groups in the least-squares procedure, the signal vanishes only in the whole-span mean.

A power spectrum (Fourier-tr of ACV) using a finite window (e.g. ± lag 2048) detects time-limited excursions of amplitude and/or phase.

We find intermodulation products at beat frequencies with the (oh so small) long-period waves (this is probably not the mechanism behind; I can try explain but it'll be lengthy).

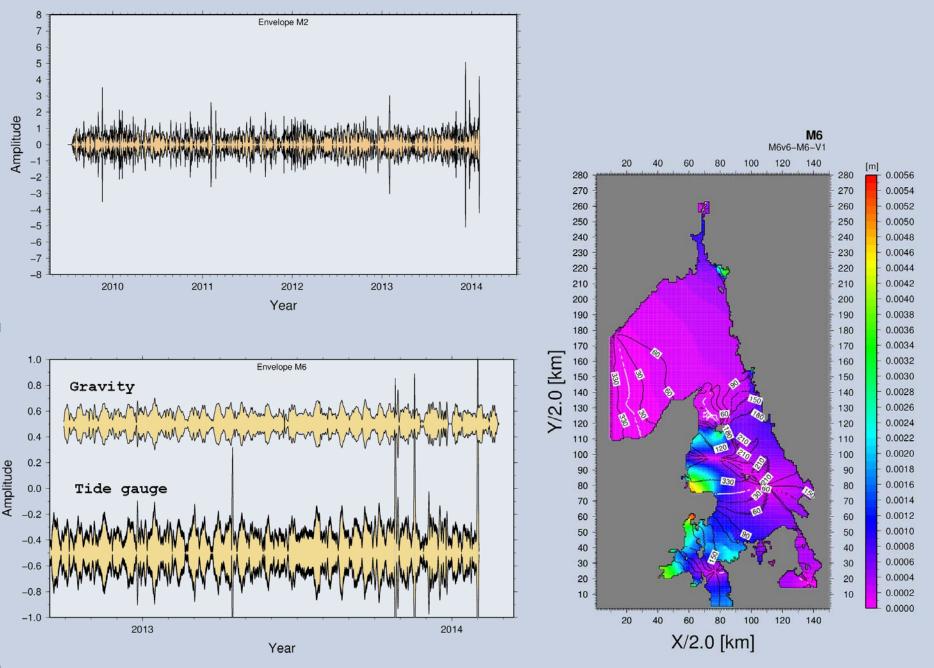


The gravimeter's M₂ "organ pipe"



M6 – band (a nonlinear tide) Shows temporal correlation of amplitudes (envelope) between tide gauge and gravity





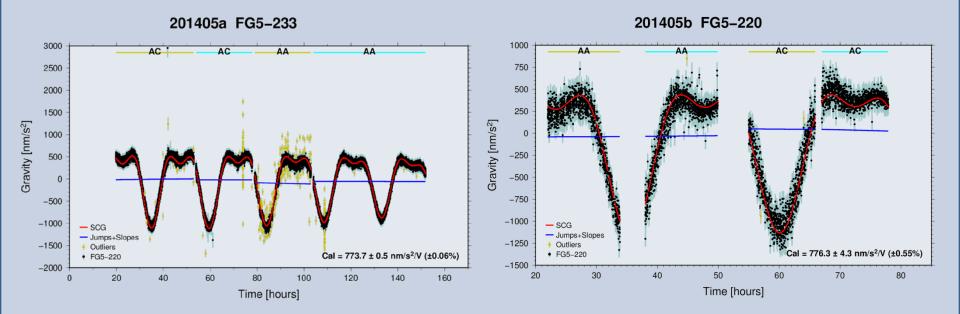


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Thus the decision to reduce AG data with SCG data, no models involved

... except the SCG drift. More precisely: no deterministic model

- Showing example from most recent campaign May 2014
- Parallel FG5x-220 (IfE) and FG5-233 (LM)
- Estimating
 - Calibration factor for SCG for whole campaign (but instruments apart)
 - Bias and slope parameters for each "project" (placement)
- Using
 - a-priori weights from AG drop files



Least-squares solution, task is **SCG-calibration** Shown:

AG – slope – bias – Acm

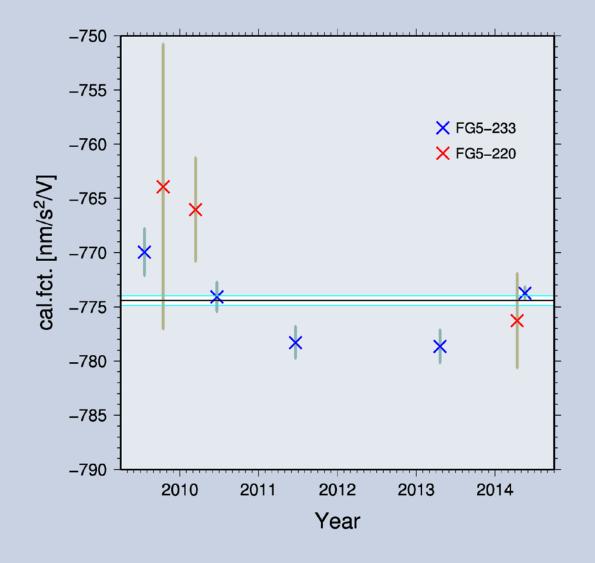
 $SCG[V] \cdot cal[nm/s^2/V] - Scm$

AG's campaign-mean

SCG's campaign-mean

Outliers as yellow error bars, accepted: light-blue

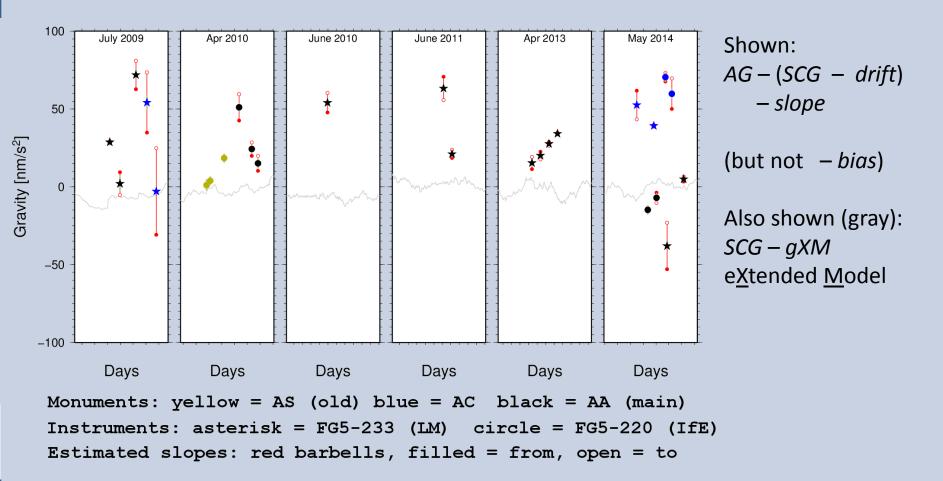
where *cal* has been determined here.



The SCG's calibration factor from AG parallel recording

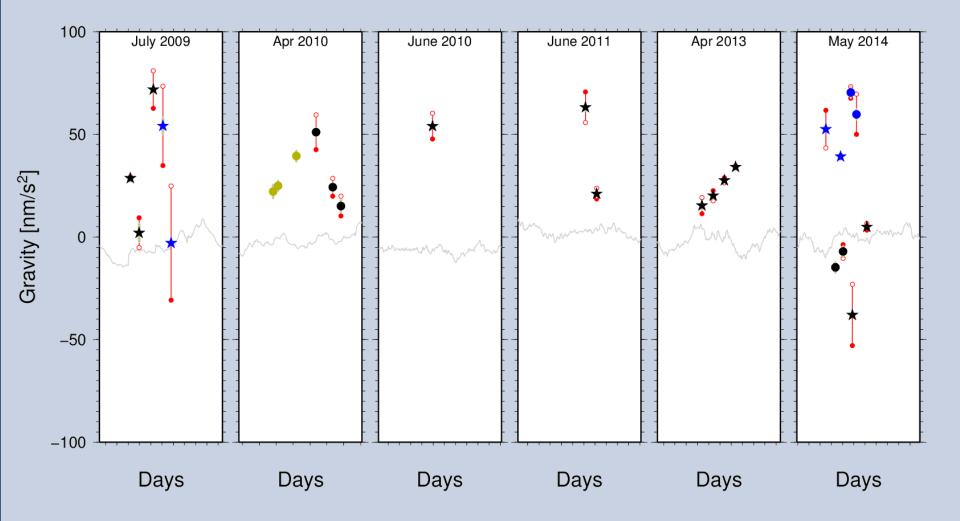
All campaigns

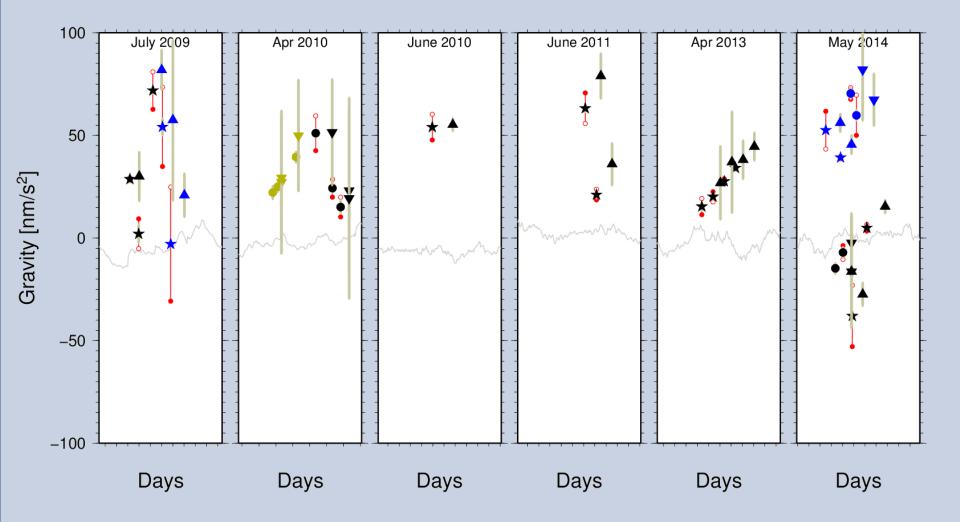
A-posteriori uncertainties (stdev) are shown



We've discussed campaign July 2009, monuments correctly identified in drop files? Project 3 on AC and 5 on AA? – well...

Scatter would not decrease if we took the start of the project as unaffected by the slopes (filled barbells).





Triangles: Weighted means, standard g-software (-like) solution

Conclusions

- We by-passed a range of models normally employed in AG observation reduction
 - Tides, atmosphere, ocean loading, pole tide
- and instead use SCG observations at the drop level
- We found drift during AG projects, slopes that occasionally exceed their standard deviations
- There are components in the SCG records which are either difficult to model (site-dependent) or have a nonstationary character. Atmosphere and annual/seasonal periodics; sea-level at temporal scale from 1 day to ~week
- Some excursions from expected values (campaigns 2010, 2011) could not be reduced. No miracle cure (there is a greater number of items in the gravimeter's error budgets than a parallel recording SCG can address).

See the poster

The Superconducting Gravimeter at Onsala Space Observatory



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http://holt.oso.chalmers.se/hgs/4me/ag2014/Poster_NKG2014_SCG-Onsala.pptx

